Precipitation extremes on multiple time scales -**Original Bartlett-Lewis model and IDF curves**

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Multi-level stochastic precipitation modeling

Precipitation processes show a high variability on spatial (localized thunderstorms) up to mesoscale hurricanes) and temporal (minutes to weeks) scales. Processes relevant on scales of extreme precipitation are not resolved in regional (RCMs) and global climate models (GCMs) and are thus not well represented. Improving their representation in these models is necessary for impact assessment and confidence in R/GCMs. Typically, parameterization of these processes is based on idealized and highly simplified models of sub-grid scale processes. Here, we explore the class of Poisson-cluster processes [Onof et al., 2000] for their suitability in subgrid-scale parametrizations. They are based on a hierarchy of Poisson processes living on different scales and are a natural approach for small scale (km, h) precipitation modelling, e.g. [Rodriguez-Iturbe et al, 1987,

Intensity-Duration-Frequency Analysis

Intensity-duration-frequency (IDF) curves show return levels as a function of rainfall duration, typically for a range of return periods. We estimate IDF curves from monthly maxima of precipitation with different aggregation times (1h to 72h) using the generalised extreme value (GEV) distribution (Coles et al., 2001) and the following relation for different durations d Koutsoyiannis et al. (1998), leading to a duration dependent formulation of the GEV[.]

$$\sigma_d = \frac{\sigma}{(d+\theta)^{\eta}} \qquad \qquad F(x;\mu,\sigma_d,\xi) = \exp\left\{-\left[1+\xi\left(\frac{x}{\sigma_d}-\mu\right)\right]^{\frac{-1}{\xi}}\right\}$$

with θ , η and σ being independent of d.

This allows simultaneous modelling of rainfall maxima across dierent durations using



Cowpertwait, 1994].

a single GEV distribution at the cost of only two additional parameters which can be estimated by maximum-likelihood (Soltyk et al., 2014).

Precipitation Model: Bartlett-Lewis rectangular pulse model cell lifetime ~ $Exp(\eta)$ Rainfall cells **Convective scale** intensity exp(1/µ_x Timescale: 1h, 3h β : expected number of cells [1/h] η : inverse of expected duration of cells [1/h] μ_x : expected intensity of cells [mm/h] Poisson Process N(t;β) of cell generation **Cell clusters** ter lifetim Synoptic scale Timescale: 12h, 24h λ : expected number of clusters [1/h] γ : inverse of expected duration of clusters [1/h] cluster generation (Poisson process N(t; λ))



IDF curves for Berlin-Dahlem

• Data: Berlin-Dahlem 5minprecipitation series time recorded by a tipping bucket from 2001 until 2013 and calculated block-maxima for every month at 1h, 3h, 6h, 12h, 24h, 48h, 72h and 96h for IDF analysis.

IDF curves for Berlin-Dahlem can be well reproduced with BLRPM simulated rainfall across the year

 Discrepancies January for due Kyrill to are (18/19.01.2007) with extreme precipitation which lead to a large shape parameter; the outstanding value does not





Visualization of July extremes as observed (RR obs , left column) and simulated by the OBL model (RR OBL, right column). Shown are short including the segments maximum observed/simulated rainfall (red vertical bars) at durations 1h (top row), 6h (middle row) and 24h (bottom) row). Additionally, the middle column shows the simulated storms (red rectangles) and (blue rectangles) cellsm corresponding to the extreme event of the simulated time series



Conclusions

- Extremes in OBL model mainly due to one long lastic cell with high intensity
- IDF relationship generally well reproduced with the OBL model, tendency to underestimate extremes with long return period
- Problems with vary rare extremes in short time-series, e.g. Kyrill in 2007

DSI 2007-07-20 14 UTC, 600hPa 1e-14 pvu^2/s

15°E

Including third moment in fitting the OBL parameters does not significantly improve the extremes in the OBL model (not shown)

Outlook: Link between stochastical simulated precipitation and DSI on multiple scales

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 $\Theta = \Theta(DSI, CAPE, ...)$

Coupling the stochastic precipitation model and a GCM requires (a) downward-coupling from the grid-scale (or larger) to the stochastic precipitation model, conceiving how the large scale flow shapes a convection permitting environment, and (b) upward-coupling, specifying the effects convection and the associated condensation with latent heat release has on the grid scale flow.



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