



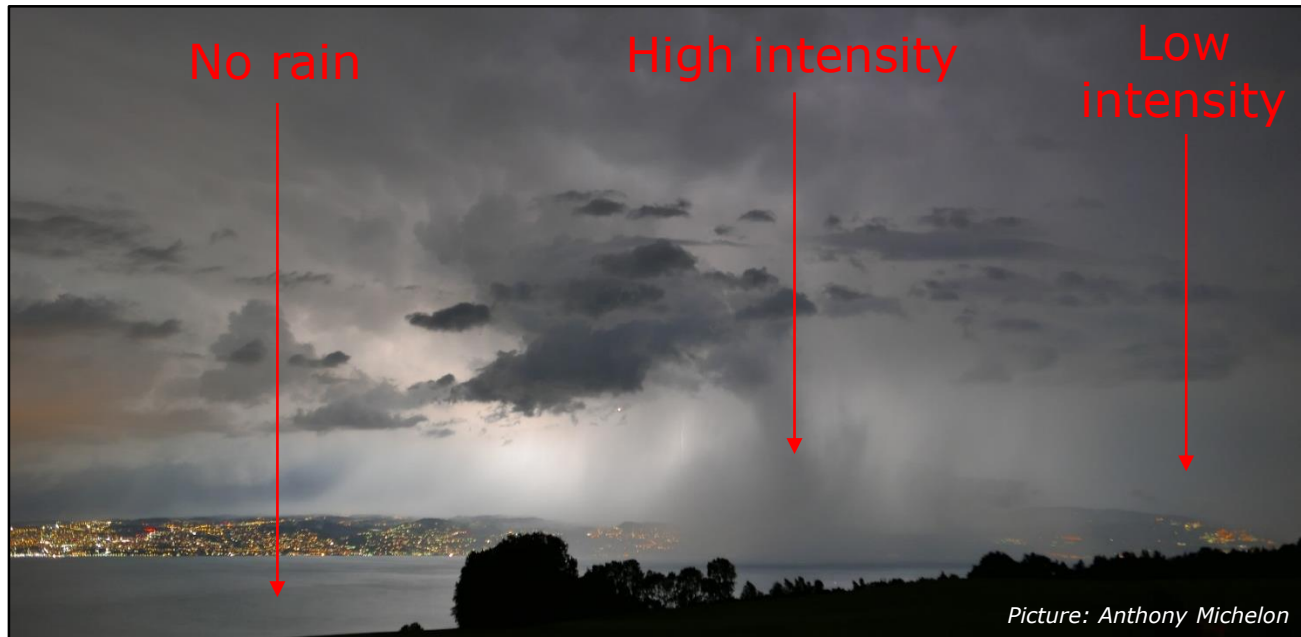
Picture: Huw Alexander Ogilvie

Sub-kilometer-scale space-time stochastic rainfall simulation

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Introduction

- When observed at the local scale, rainfall appears highly variable in space and in time within rain events.

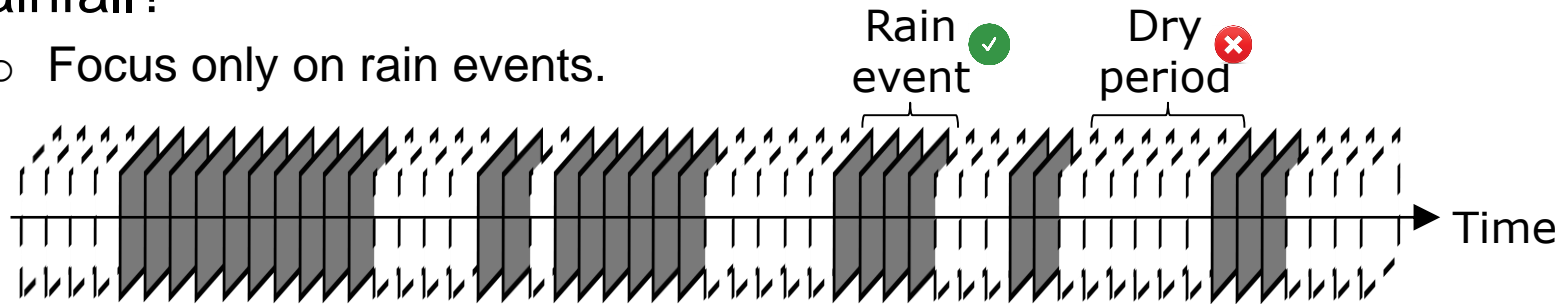


=> How can we measure and reproduce the statistical behavior of rainfall?

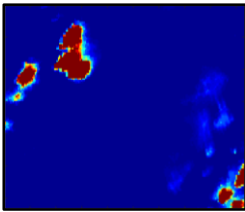
Introduction

- How can we measure and reproduce the statistical behavior of rainfall?

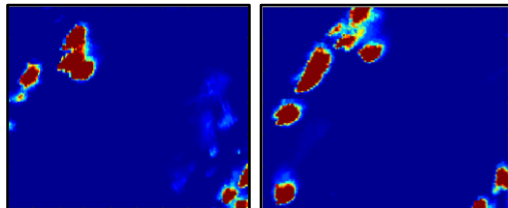
- Focus only on rain events.



- Within rain events, pay a particular attention to space-time dependencies:



- Spatial correlation?
- Spatial patterns?
- Nature of rain / no-rain transitions?



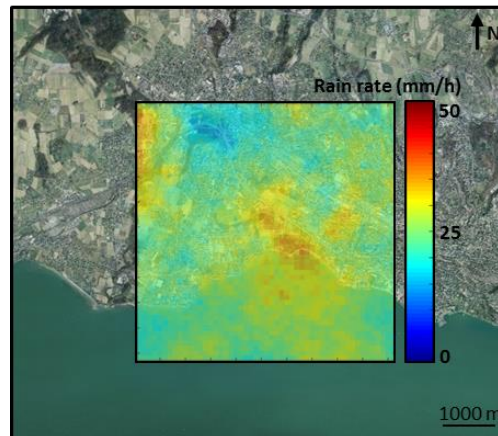
- Temporal correlation?
- Morphing of spatial patterns?
- Advection of rain storm?

Introduction

- How can we measure and reproduce the statistical behavior of rainfall?

=> Use a stochastic rainfall model.

- Input data: rain rate time series (high resolution rain gauges).
- Parameter inference:
 - Calibration of the model.
 - The calibrated model gives insights on the structure of rainfall.
- Simulation: generate synthetic rain fields which reproduce the structure of observed rain fields.



Instruments

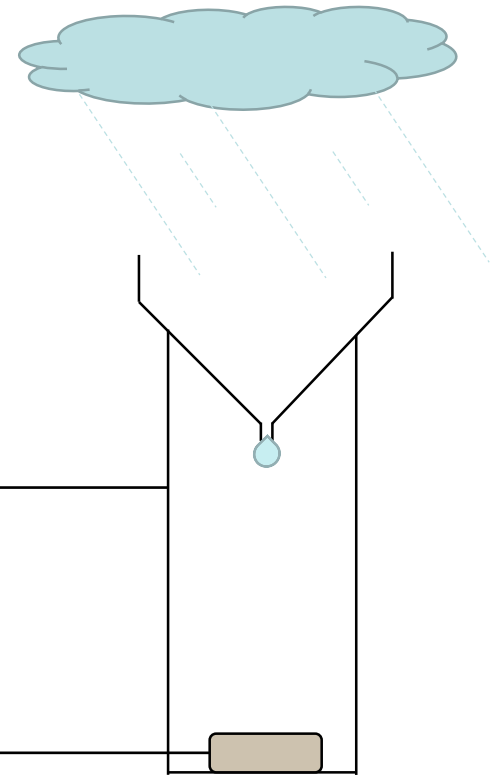
- Questions:
- How to measure local rain fields?
 - Which data to feed a local scale stochastic rainfall model?



Instruments

Drop counting rain gauge

- Drop counting rain gauge (Pluvimate).
- Operation principle: counts calibrated drops instead of bucket tips.

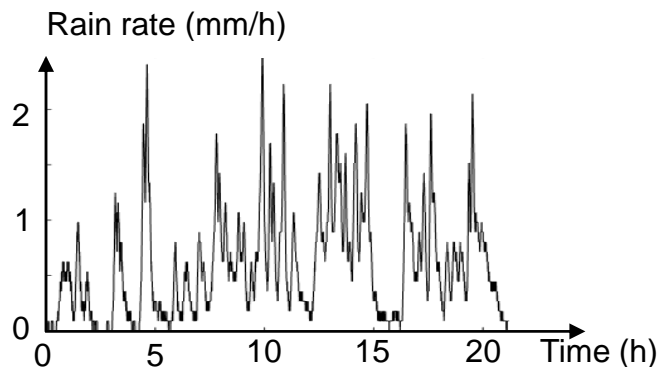
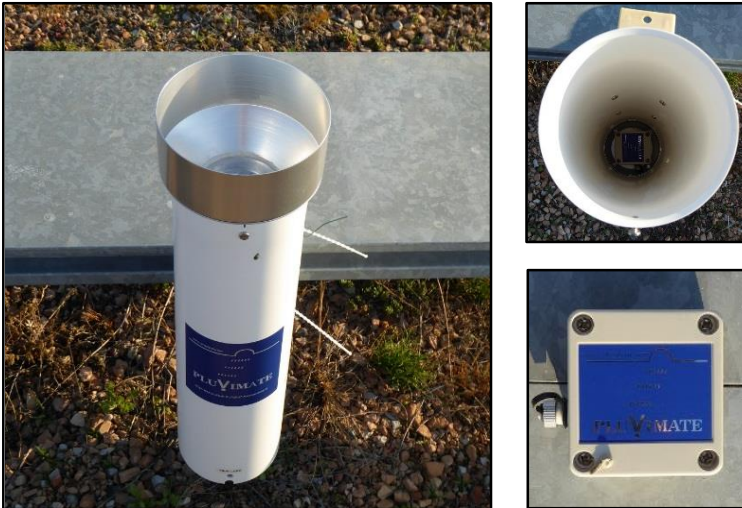


Instruments

High resolution

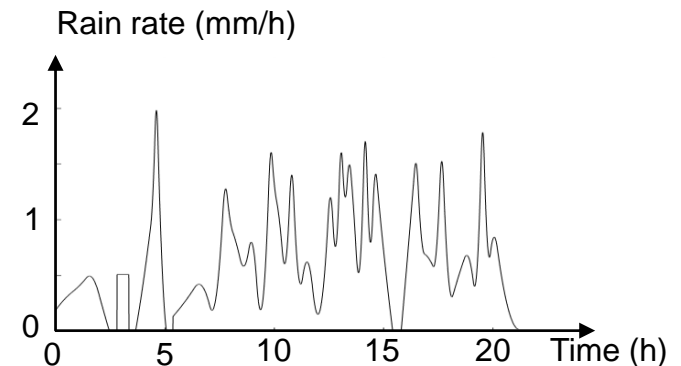
Pluvimate

- Operation principle: Drop counting.
- Rain height resolution: 0.01mm.
- Sampling rate: 30sec – 1min.



Watchdog 1120

- Operation principle: Tipping bucket.
- Rain height resolution: 0.1-0.25mm.
- Sampling rate: 5 – 10 min.

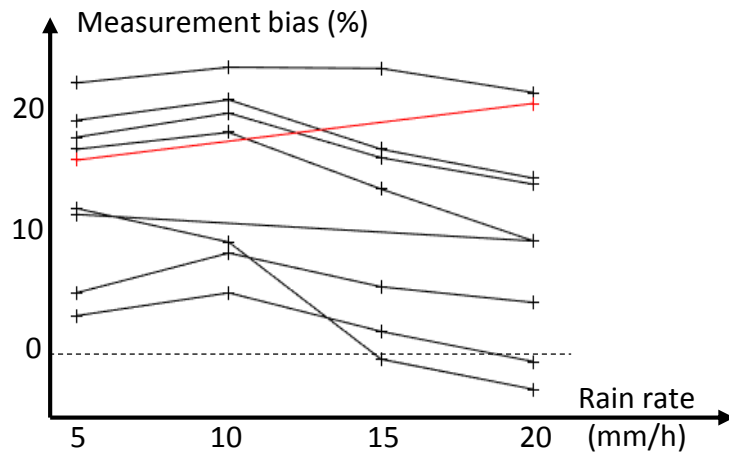


Instruments

Calibration

Absolute calibration

Artificial rain, 0-20mm/h

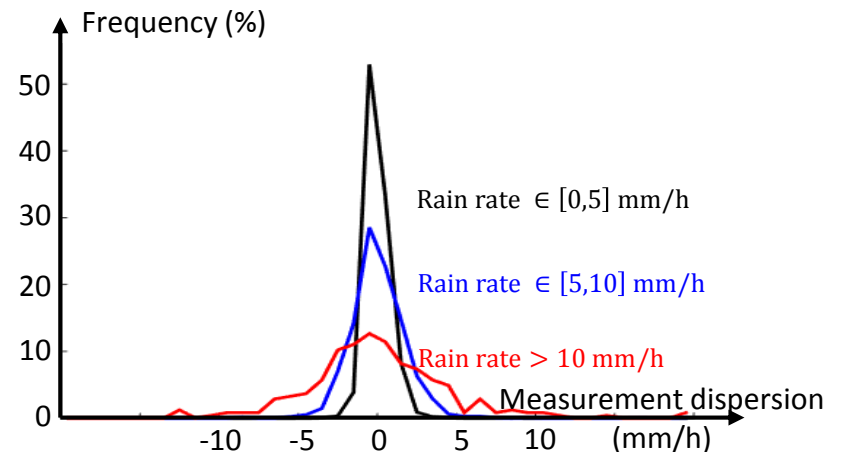


Black: Pluviate device

Red: Tipping bucket rain gauge (for comparison)

Relative calibration

Natural rain, 0-40mm/h



Rain rate $\in [0,5]$ mm/h

Rain rate $\in [5,10]$ mm/h

Rain rate > 10 mm/h

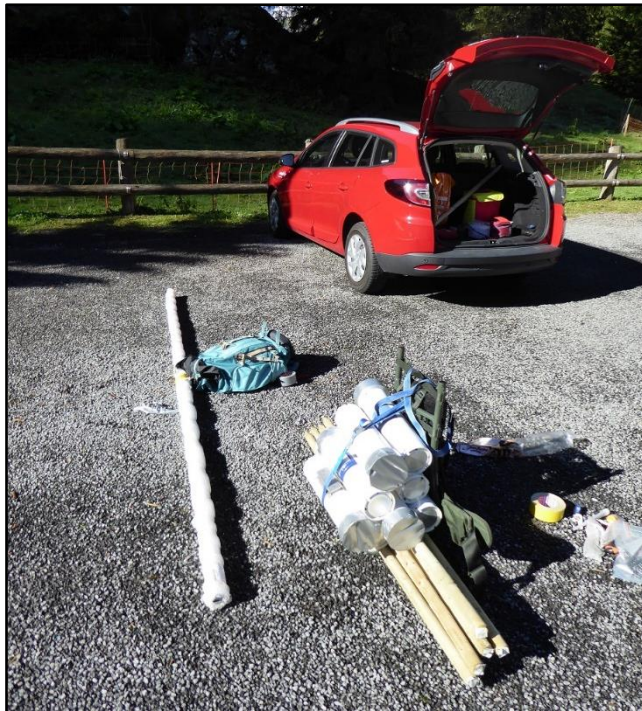
Measurement dispersion

Instruments

Low cost and easy to set up

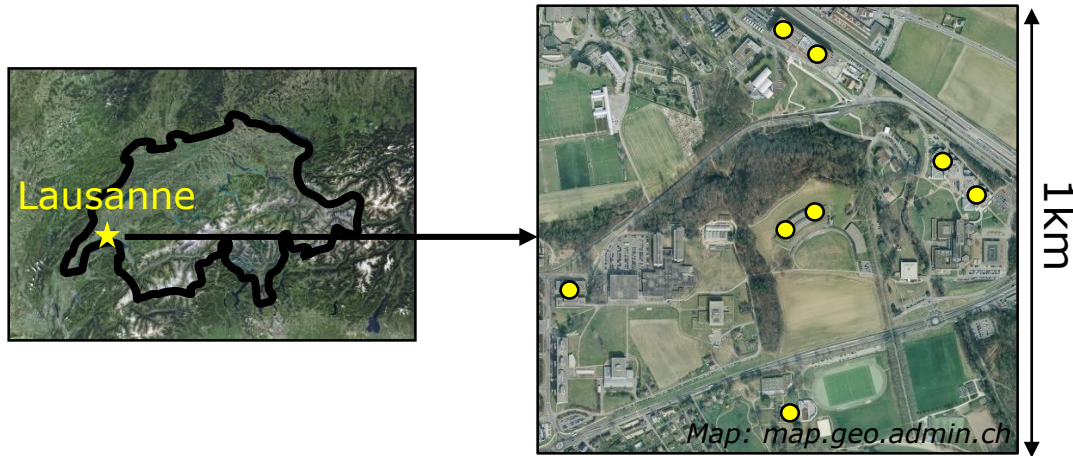
- Relatively cheap (~500\$ per device).
- Easy to set up (light, no moving parts, low power consumption).
- Few maintenance (except low storage capacity).

=> Dense networks with many gauges, even in mountains.



Instruments

Experimental network

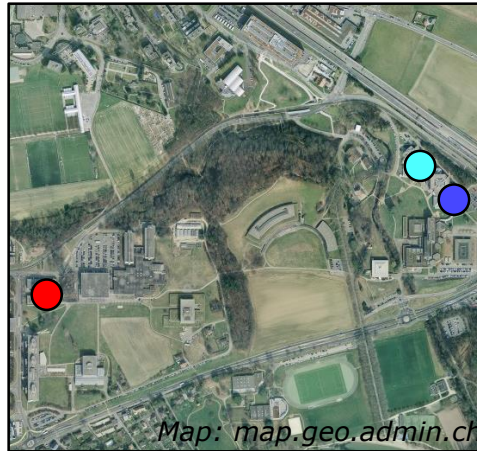


- 8 rain gauges.
- 1km x 1km.
- Sampling rate: 30sec.

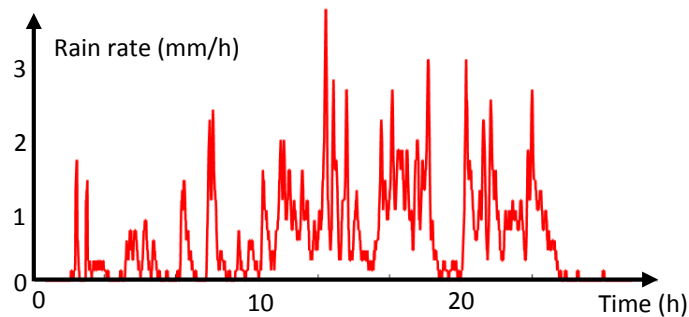


Instruments

Observed rainfall structure

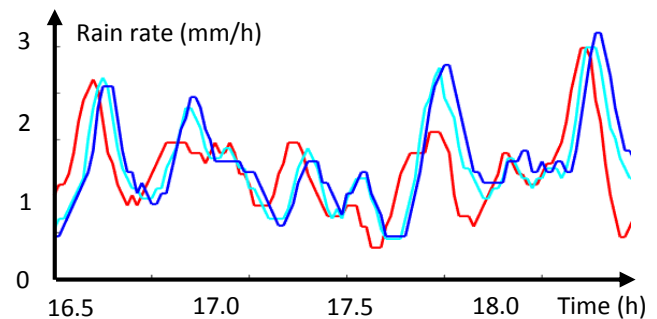


Single site observations



- Mass of zero measurements.
- Smooth dry/wet transitions.

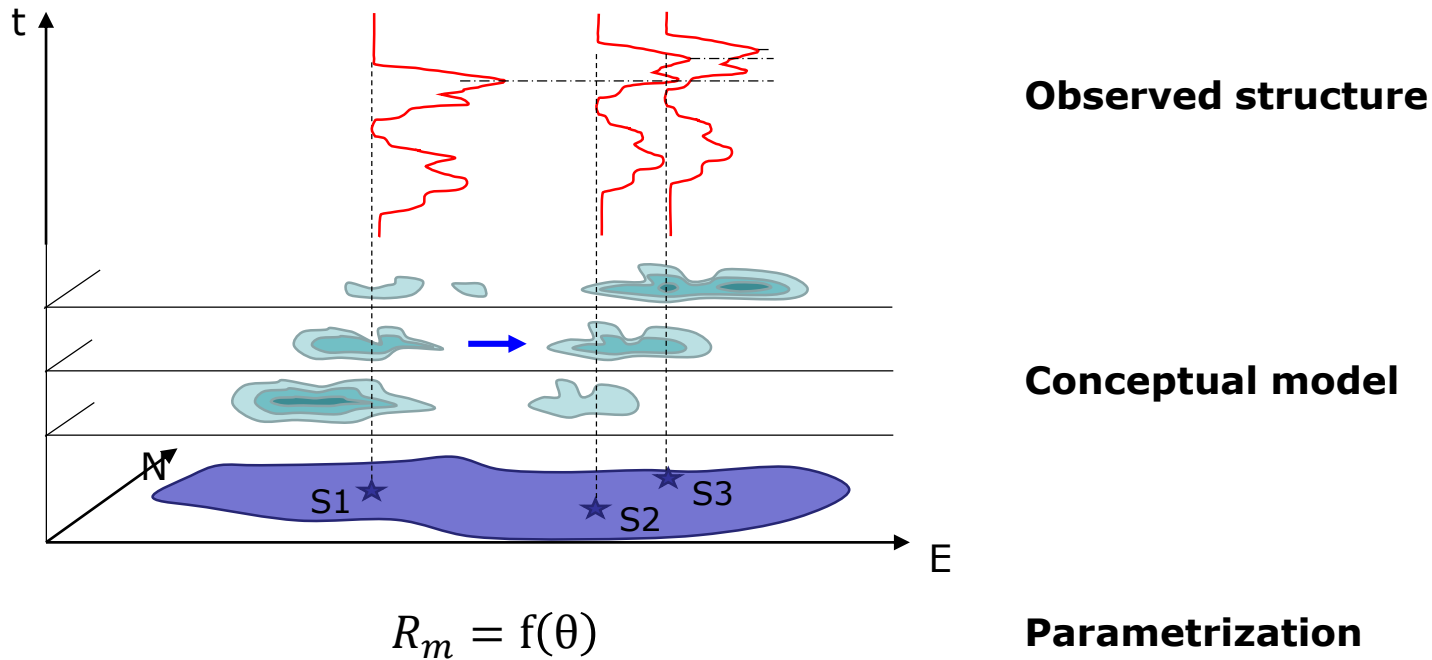
Multi-sites observations



- Spatial variability of time series.
- More similarities for close gauges.
- Time shift for distant gauges.

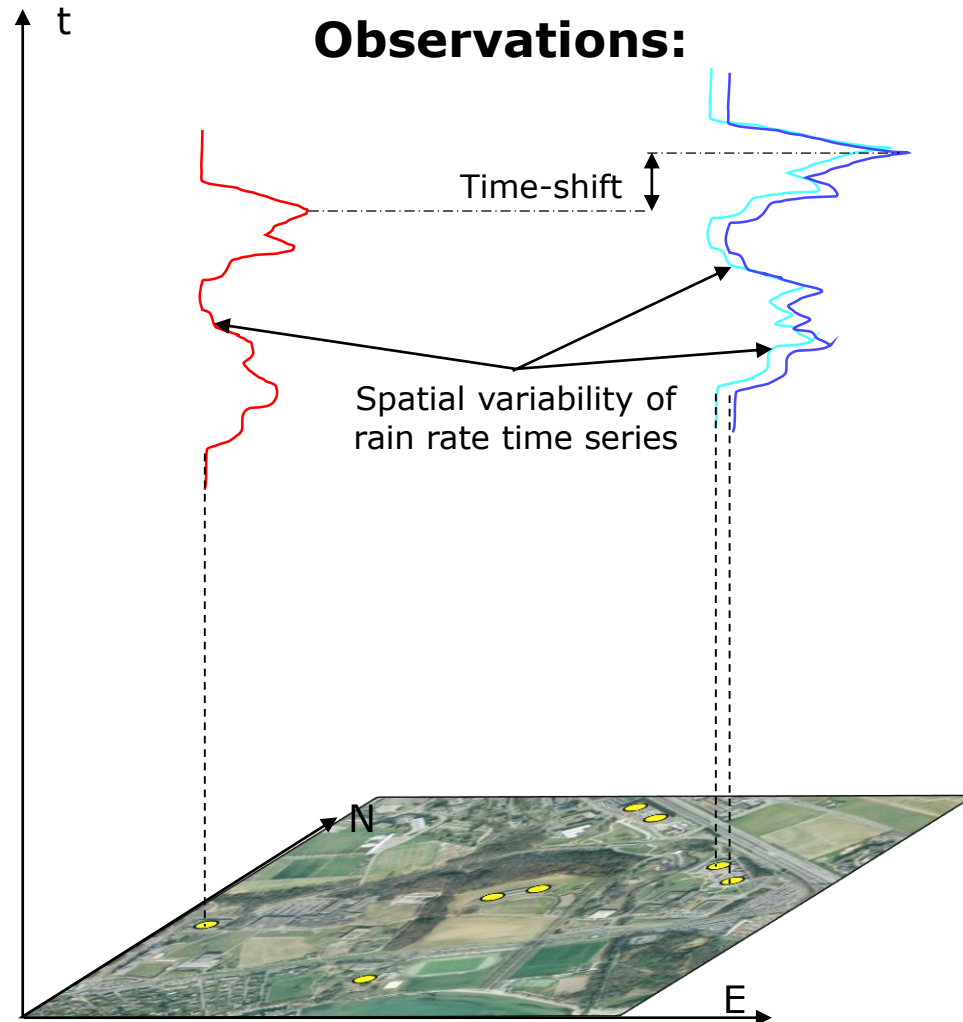
Stochastic rainfall model

Question: How to choose a stochastic model which can handle the features of rainfall arising from Pluvimate data?



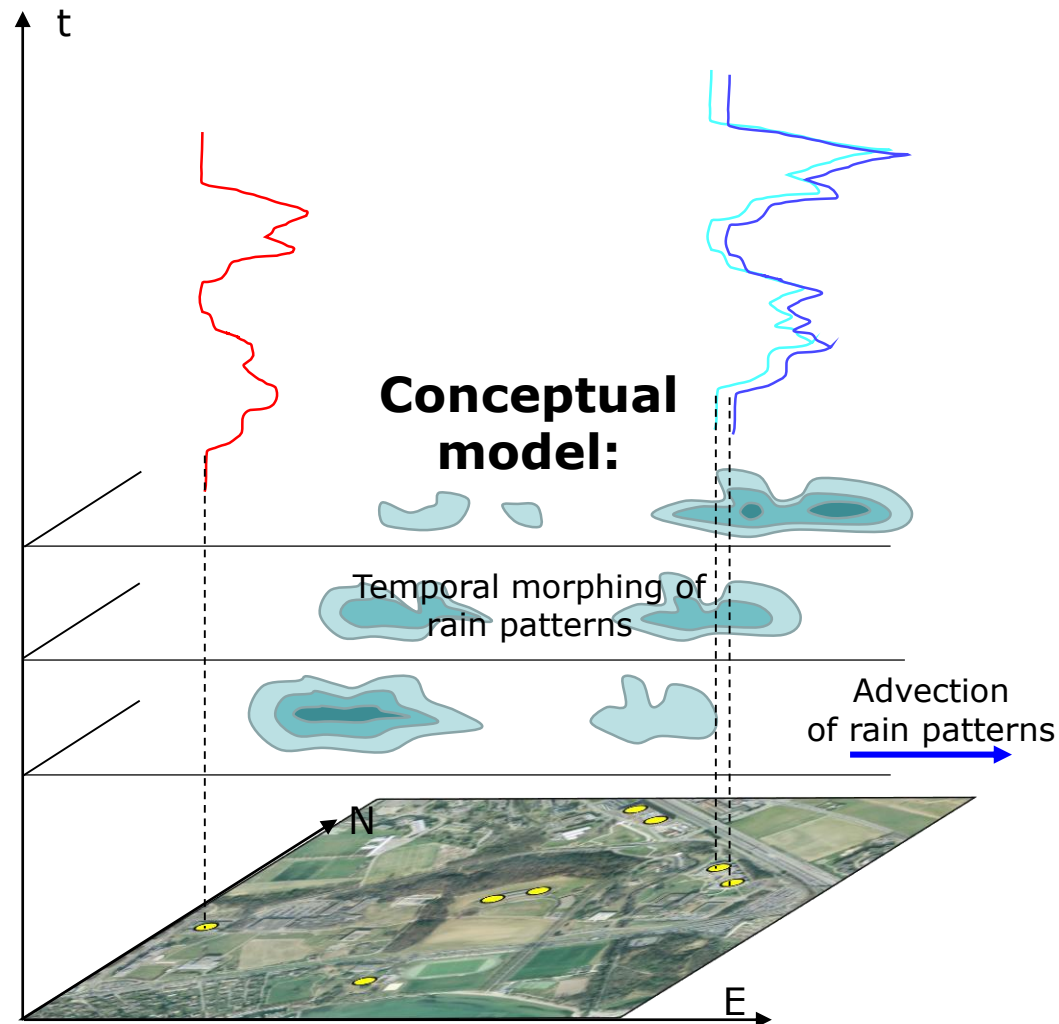
Stochastic rainfall model

Conceptual model



Stochastic rainfall model

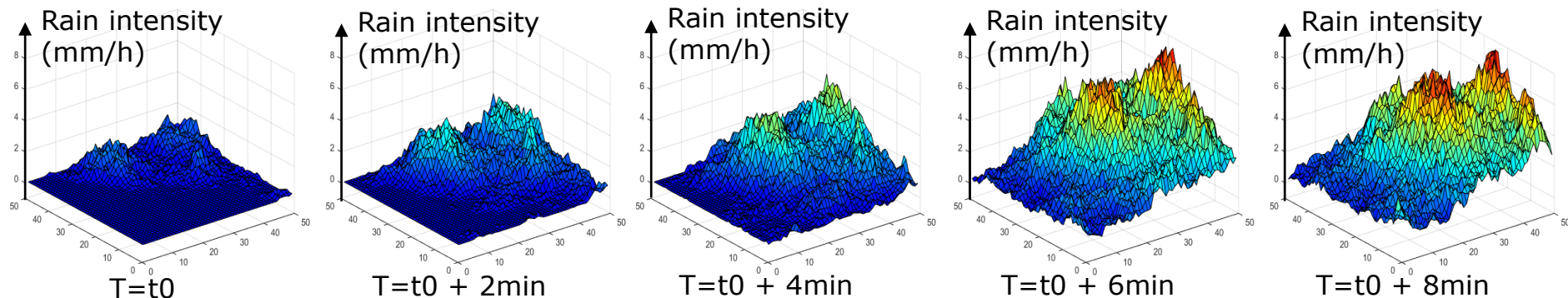
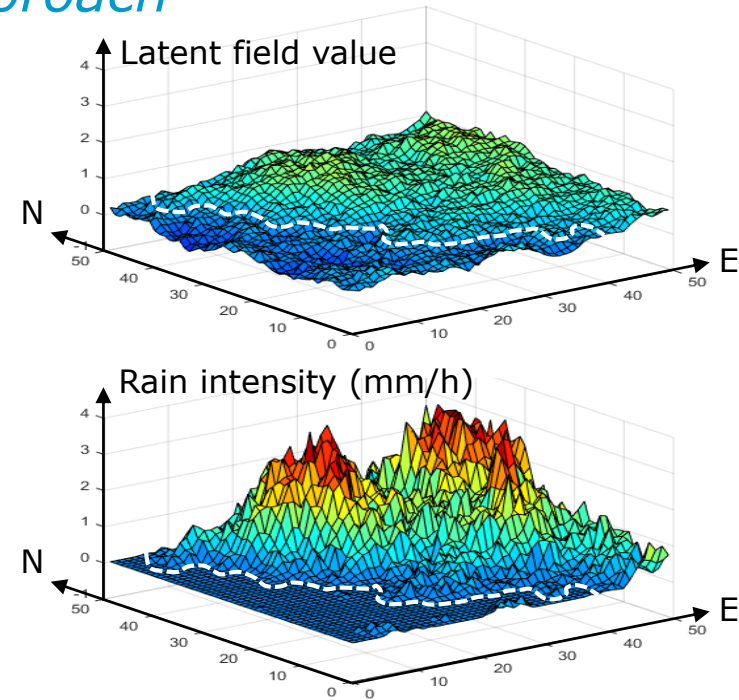
Conceptual model



Stochastic rainfall model

Overall modelling approach

- Latent Gaussian random field.
- Precipitation arise from the latent field through truncation and transformation.
- Rainfall dynamics is modeled by an asymmetric and non-separable covariance function.



Stochastic rainfall model

Parametrization

- Latent Gaussian random field:
 - Spatio-temporal coordinates: (\mathbf{s}, t) .
 - Standardized and stationary multivariate Gaussian random field: $Y(\mathbf{s}, t)$.
- Precipitation arise from the latent field through truncation and transformation:
$$\begin{cases} R_m(\mathbf{s}, t) = 0 & \text{if } Y(\mathbf{s}, t) \leq a_0 \\ R_m(\mathbf{s}, t) = \left(\frac{Y(\mathbf{s}, t) - a_0}{a_1} \right)^{\frac{1}{a_2}} & \text{if } Y(\mathbf{s}, t) > a_0 \end{cases}$$
- Rainfall dynamics is modeled by an asymmetric and non-separable covariance function $\rho(\mathbf{s}, t)$:
 - Advection is modeled by a single vector V : $\rho(d\mathbf{s} - V \cdot dt, dt) = \rho_L(d\mathbf{s}, dt)$.
 - Diffusion / morphing is modelled by a non-separable covariance in a Lagrangian reference frame:

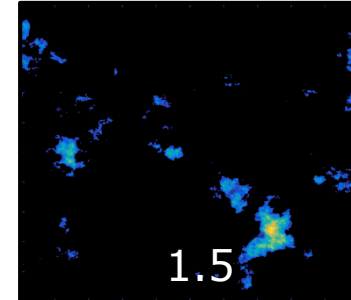
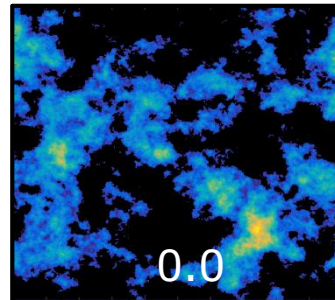
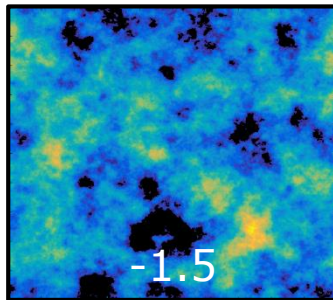
$$\rho_L(d\mathbf{s}, dt) = \frac{1}{\left(\left(\frac{dt}{d} \right)^{2\delta} + 1 \right)} \cdot \exp \left(- \frac{\left(\frac{d\mathbf{s}}{c} \right)^{2\gamma}}{\left(\left(\frac{dt}{d} \right)^{2\delta} + 1 \right)^{\beta\gamma}} \right)$$

Stochastic rainfall model

Parametrization

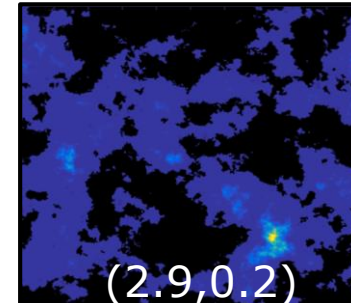
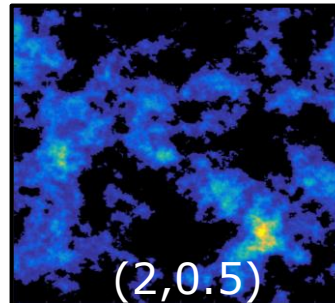
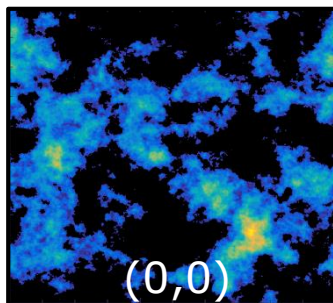
- Transform function:

- Truncation => proportion of dry areas: $R_m(\mathbf{s}, t) = 0$ if $Y(\mathbf{s}, t) \leq a_0$



- Transform function => skewness of the marginal distribution:

$$R_m(\mathbf{s}, t) = \left(\frac{Y(\mathbf{s}, t) - a_0}{a_1} \right)^{1/a_2} \text{ if } Y(\mathbf{s}, t) > a_0$$



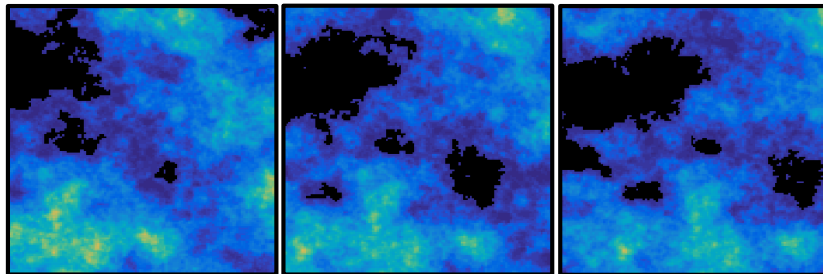
Stochastic rainfall model

Parametrization

- Covariance of the latent field:

- Advection vector: $\rho(ds - \mathbf{V}.dt, dt) = \rho_L(ds, dt)$.

V=1

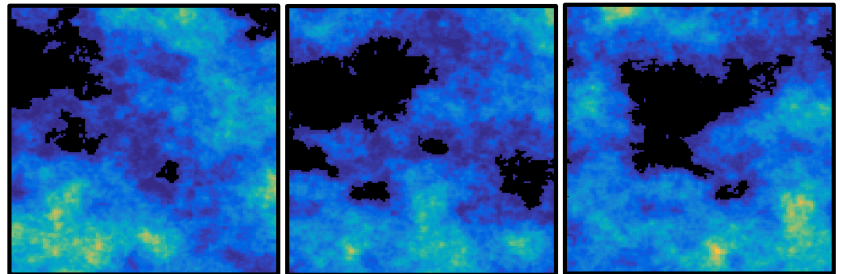


t=1

t=2

t=3

V=2



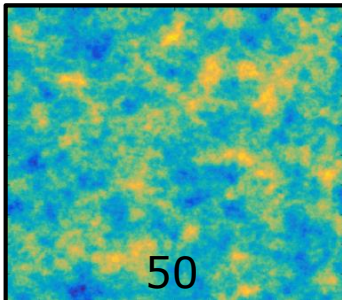
t=1

t=2

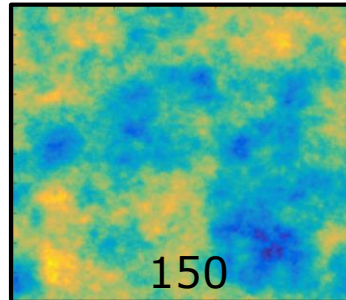
t=3

- Spatial dependencies: $\rho_L(ds, dt) = \frac{1}{\left(\left(\frac{dt}{d}\right)^{2\delta} + 1\right)} \cdot \exp\left(-\frac{\left(\frac{ds}{c}\right)^{2\gamma}}{\left(\left(\frac{dt}{d}\right)^{2\delta} + 1\right)^{\beta\gamma}}\right)$

Range (parameter c): size of patterns

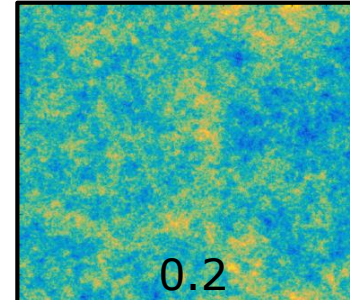


50

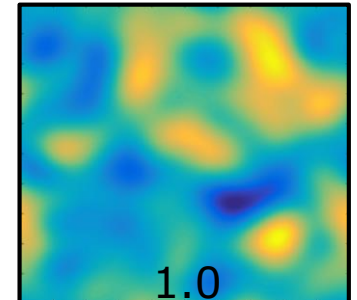


150

Shape (parameter γ): shape of patterns



0.2



1.0

Stochastic rainfall model

Parameter inference

- The inference of model parameters allows to:
 - Calibrate the model.
 - Gain insights on the structure of rainfall.
- A Bayesian approach is selected to account for the uncertainty on model parameters.
 - Use a Metropolis-Hastings sampler, which requires:
 - A statistical model to calibrate => stochastic rainfall model.
 - A calibration dataset => Rain rate time series.
 - The likelihood of observations given model parameters => Easy to derive thanks to the assumption of multivariate Gaussian latent field.

Metropolis-Hasting algorithm

- 1) Initialize model parameter $\theta \in \Theta$
- 2) (a) Generate $\theta^* \sim q(\theta^*|\theta)$ and $u \sim U_{[0,1]}$ (q = proposition kernel)
(b) If $u < \min\left(1, \frac{l(R_m|\theta^*) \times q(\theta|\theta^*)}{l(R_m|\theta) \times q(\theta^*|\theta)}\right)$, then $\theta = \theta^*$
- 3) Iterate 2)

Stochastic rainfall model

Parameter inference

- For large datasets, the calculation of the full-likelihood is computationally infeasible.

$$L(R_m|\theta) = -0.5[\log|\Sigma_{++}| - Z_+^t \cdot \Sigma_{++}^{-1} \cdot Z_+ - N_+ \cdot \log(2\pi)] \longrightarrow |\cdot| \text{ and } \cdot^{-1} \text{ for large matrices}$$
$$+ \log \phi_{N_0}(a_0|\Sigma_{+0}\Sigma_{++}^{-1}, \Sigma_{00} + \Sigma_{+0}\Sigma_{++}^{-1}\Sigma_{0+}) \longrightarrow \text{joint cdf } \phi_{N_0} \text{ at many points}$$

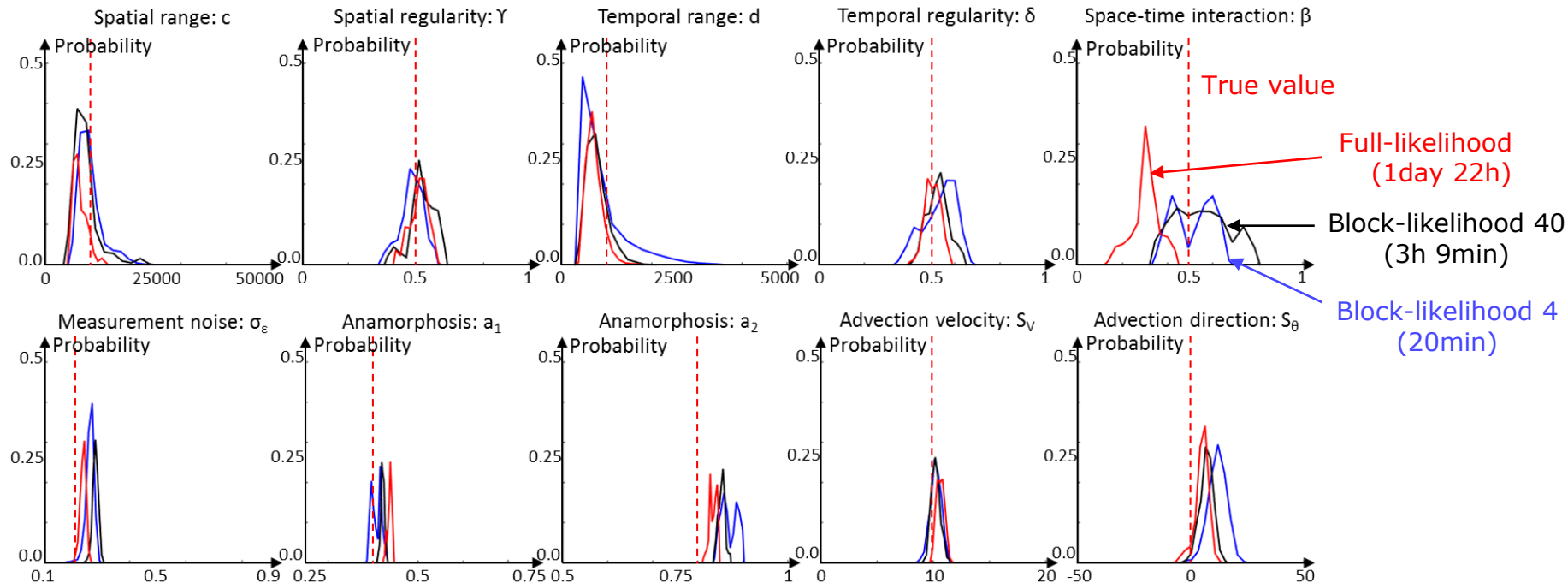
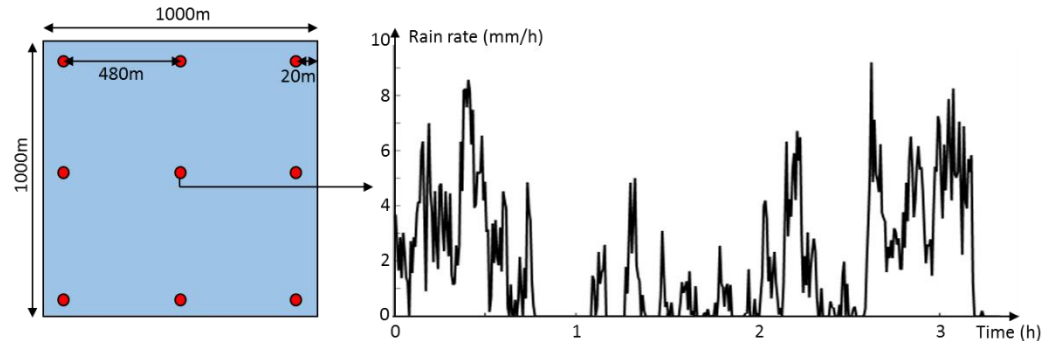
- Approximations are required:
 - For positive measurements, the likelihood is evaluated for small blocks.
 - Blockwise likelihood: $l(\theta|R_{I_+}) \approx \prod_{p=1}^{N_\tau} l(\theta|R_{BI_+})$
 - For zero measurements, the censored values of the latent field are simulated by a Gibbs sampler within the Metropolis-Hasting algorithm.

Stochastic rainfall model

Parameter inference

- In practice, the previous approximations don't worsen the inference of parameters:

- Test on a synthetic case:



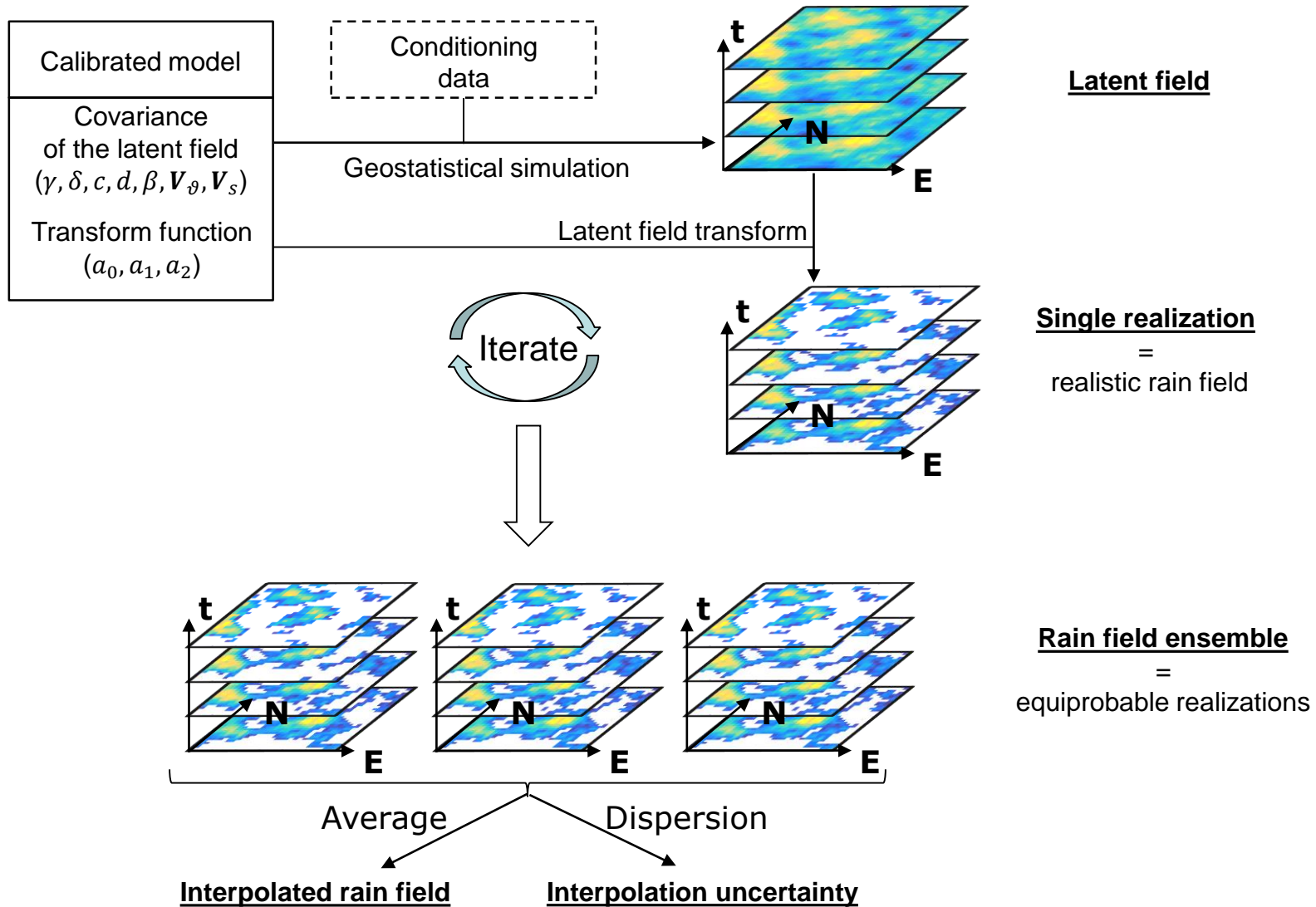
Stochastic rainfall model

Simulation

- The posterior distribution of the stochastic rainfall model can be used to generate synthetic rain fields, which are useful to:
 - Simulate rainfall over the space-time domain of interest.
 - Interpolate rain at any ungauged location (or time step):
 - Predicted value.
 - Assessment of prediction error.
- Simulation method:
 - The latent field is first obtained by geostatistical simulation.
 - Synthetic rain fields are derived by censoring & transforming the latent field.

Stochastic rainfall model

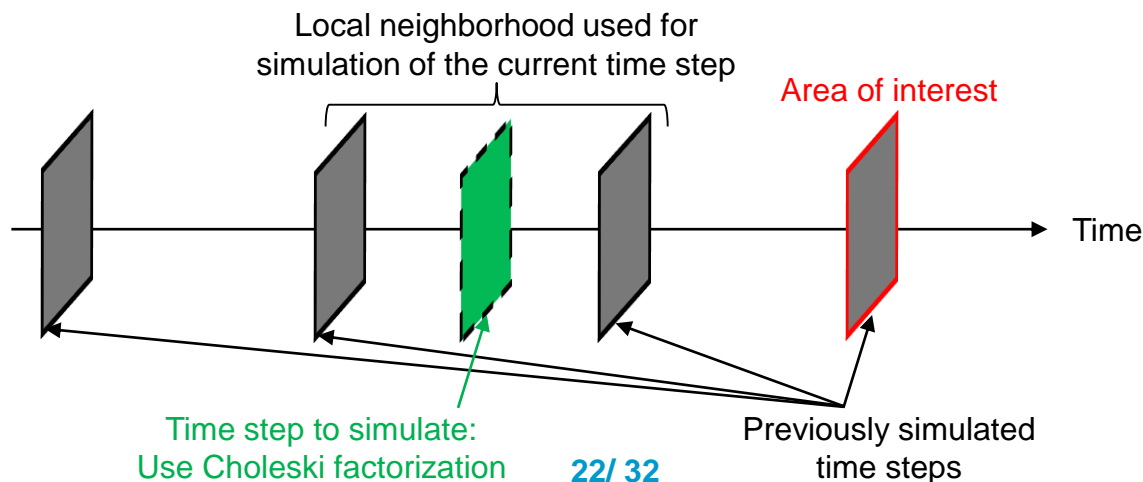
Simulation



Stochastic rainfall model

Simulation

- For large space-time simulation grids, exact geostatistical simulations of the latent field are computationally unfeasible.
 - Require Choleski factorization of large covariance matrices.
- Usual fast simulation methods (turning bands, FFT-based, etc.) cannot be easily applied in the current context.
- An ad-hoc simulation method has been developed:
 - Choleski factorization in the space dimension.
 - Multigrid Sequential Gaussian Simulation (SGS) in the time dimension.

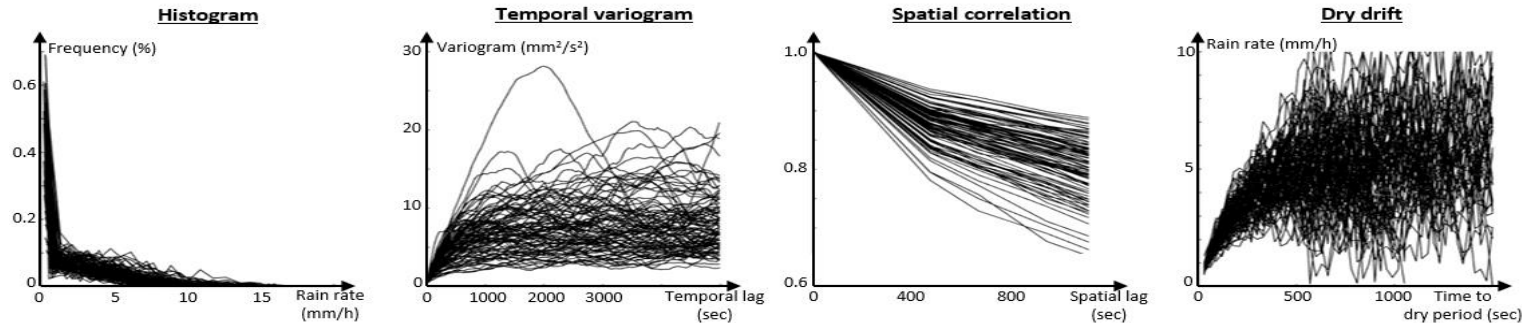


Stochastic rainfall model

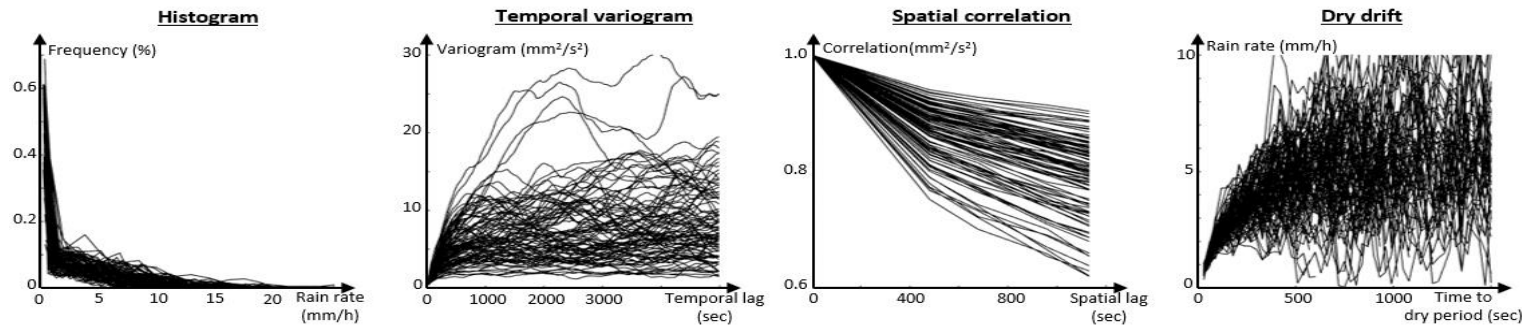
Simulation

- In practice, the proposed simulation method does not worsen the reproduction of rainfall statistics.

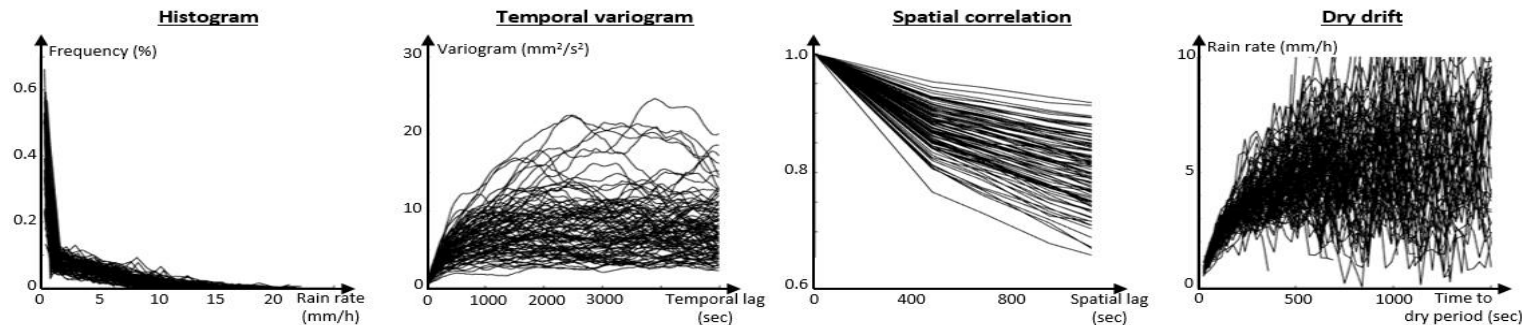
Exact simulation
(Choleski decomposition)



Approximated simulation
(40 neighboring time steps)

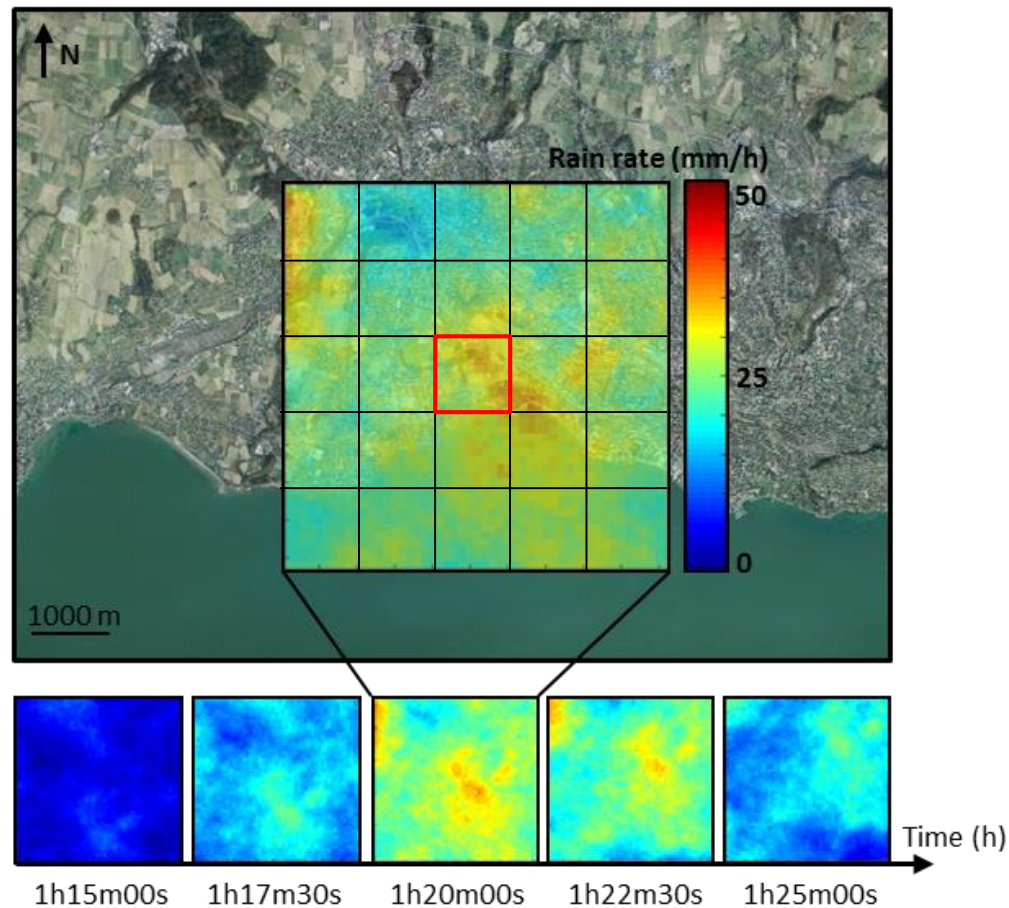


Approximated simulation
(4 neighboring time steps)



Application

Question: What happens within a single radar pixel?

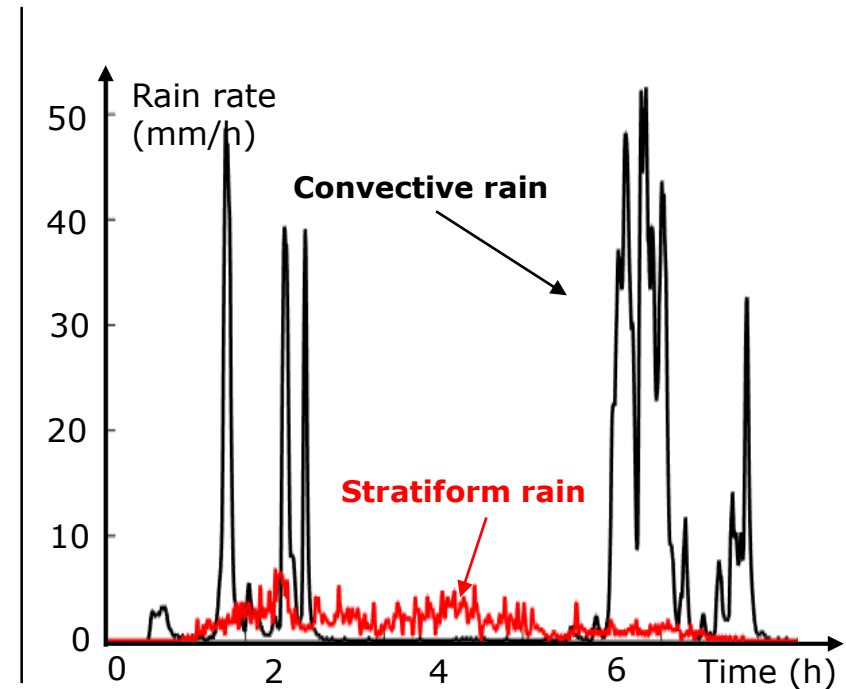
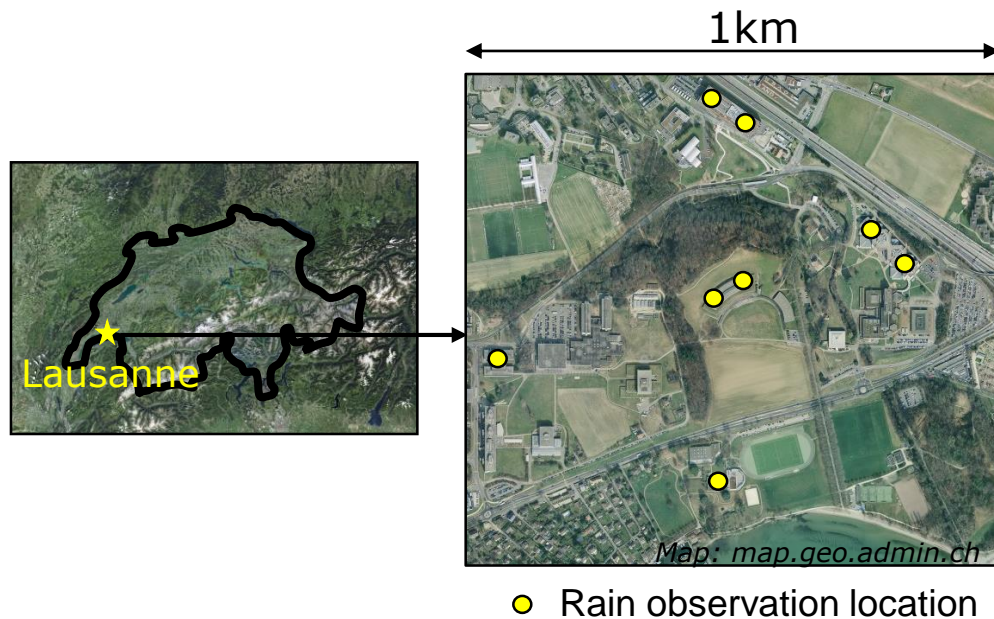


Application

Experimental setup

- The proposed model is applied to two rain events in order to:
 - Assess the space-time structure of the rain events (parameter inference step).
 - Generate synthetic rain fields conditioned to observations (simulation step).
- Experimental setup:

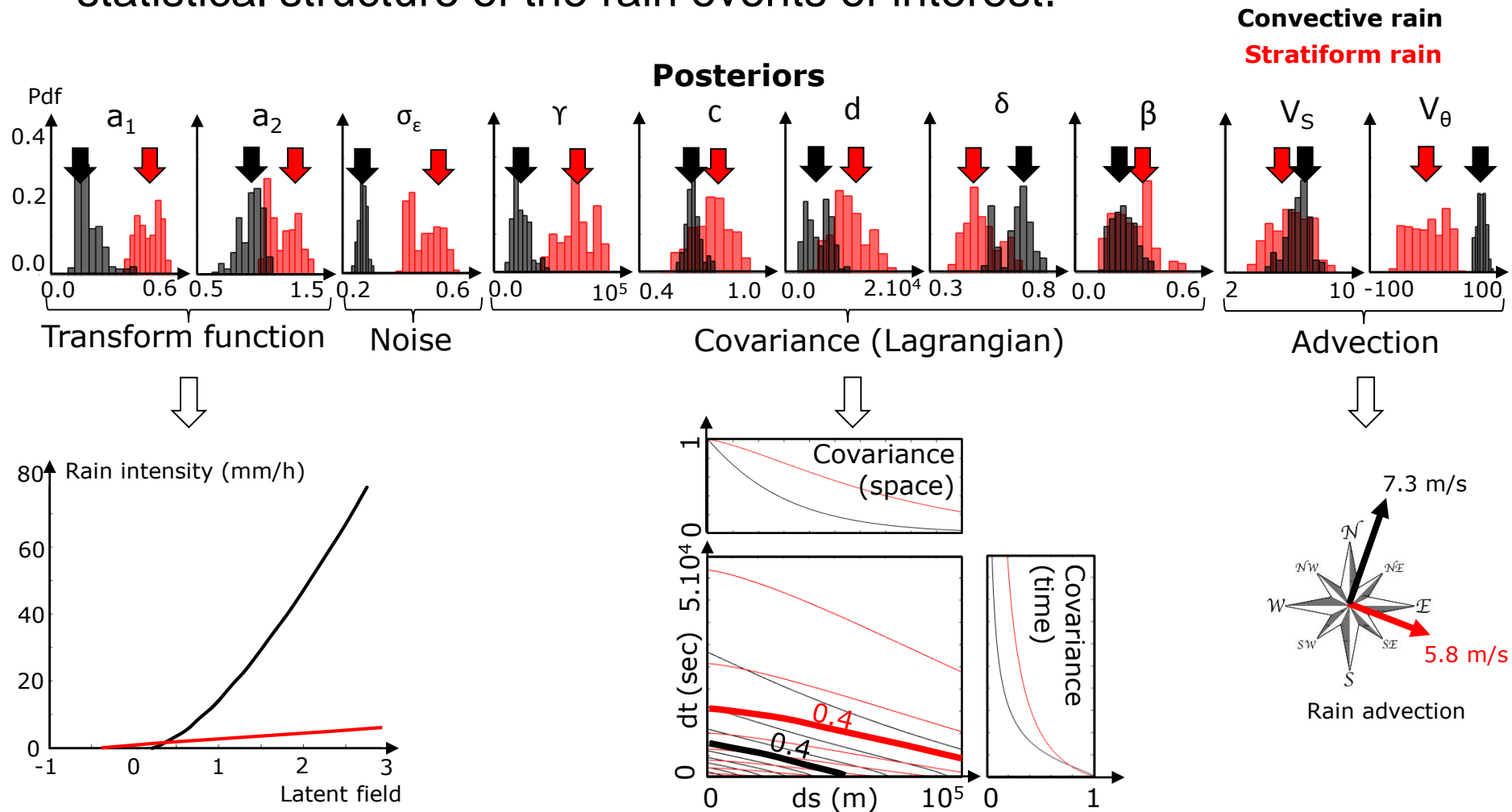
- Rain events of interest:



Application

Parameter inference

- Model parameters are inferred to assess the space – time – intensity statistical structure of the rain events of interest.

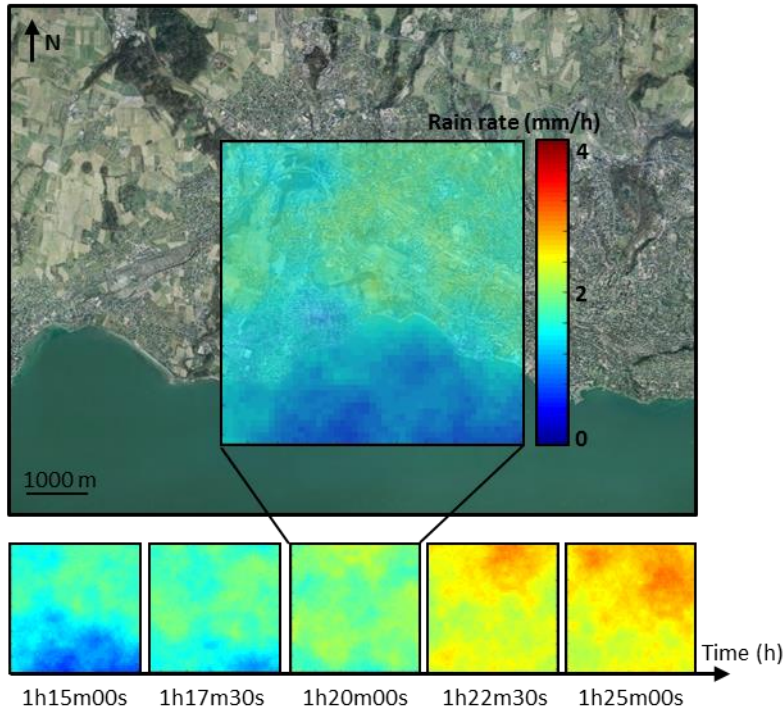


Application

Simulation of synthetic rain fields

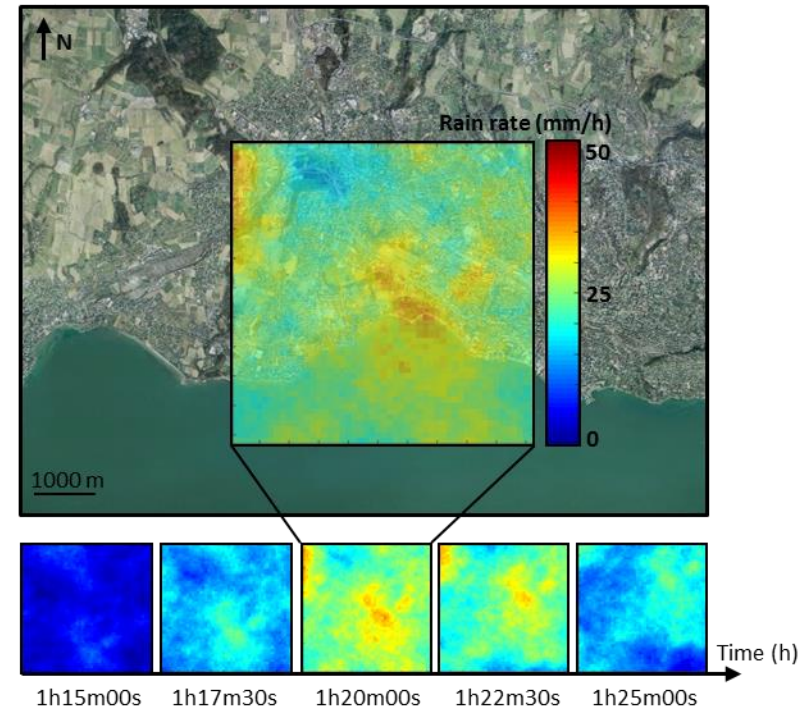
- The inferred model parameters lead to different simulated rain fields.

Stratiform rain



- Gentle rain, low intermittency.
- Low space-time variability.
- Eastward advection.

Convective rain



- Heavy rain, intermittent.
- High space-time variability.
- Northward advection.

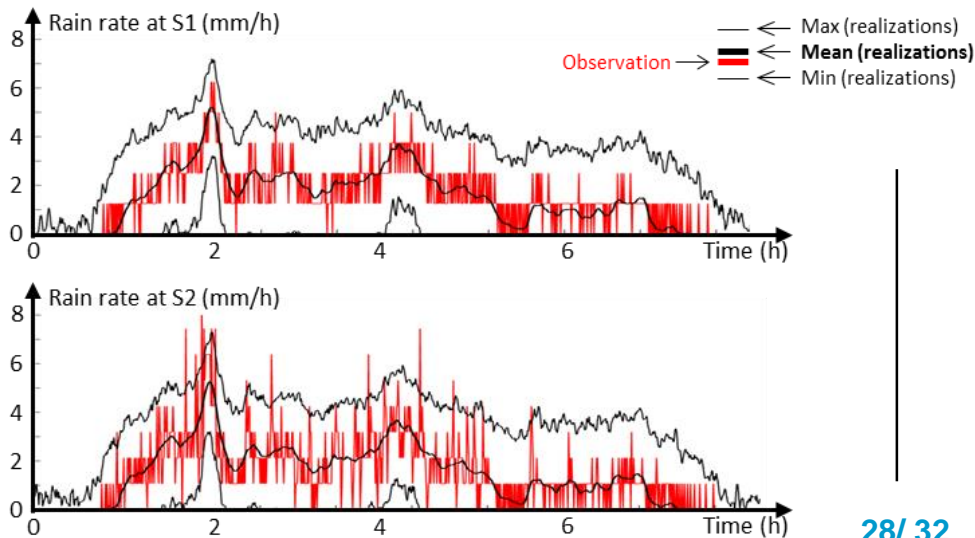
Application

Rain prediction

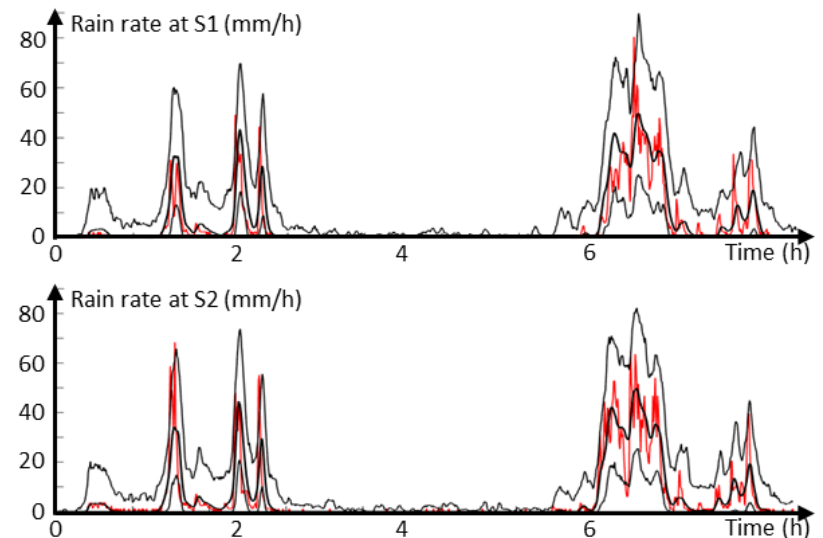
- An ensemble of 100 realizations can be used to predict rainfall at ungauged locations and to assess the prediction uncertainty.



Stratiform rain



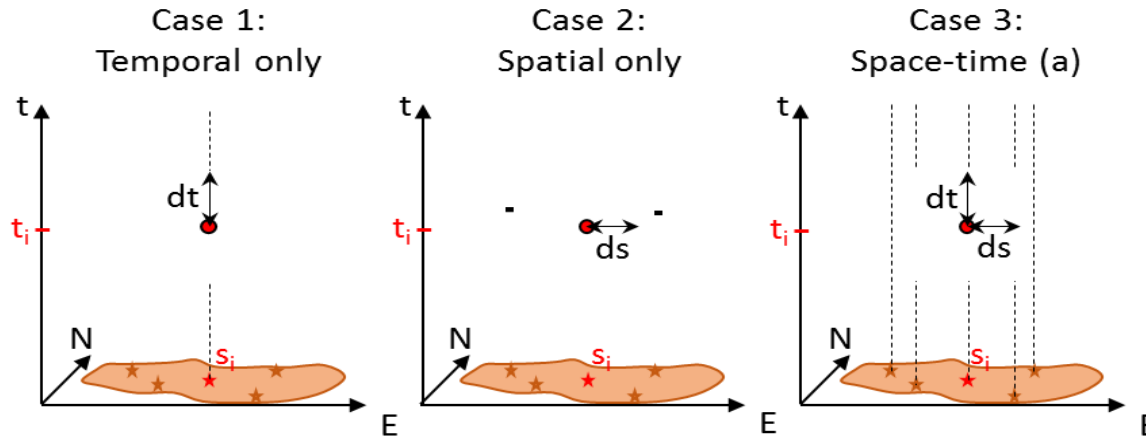
Convective rain



Application

Cross-validation

- Cross-validation to assess the reliability of rainfall prediction:



Stratiform rain

	C1	C2	C3
dt = 30s ds = 50m	MAE = 0.70 Bias = -0.01	MAE = 0.69 Bias = -0.001	MAE = 0.71 Bias = -0.06
dt = 300s ds = 500m	MAE = 0.74 Bias = -0.01	MAE = 0.70 Bias = -0.002	MAE = 0.75 Bias = -0.006

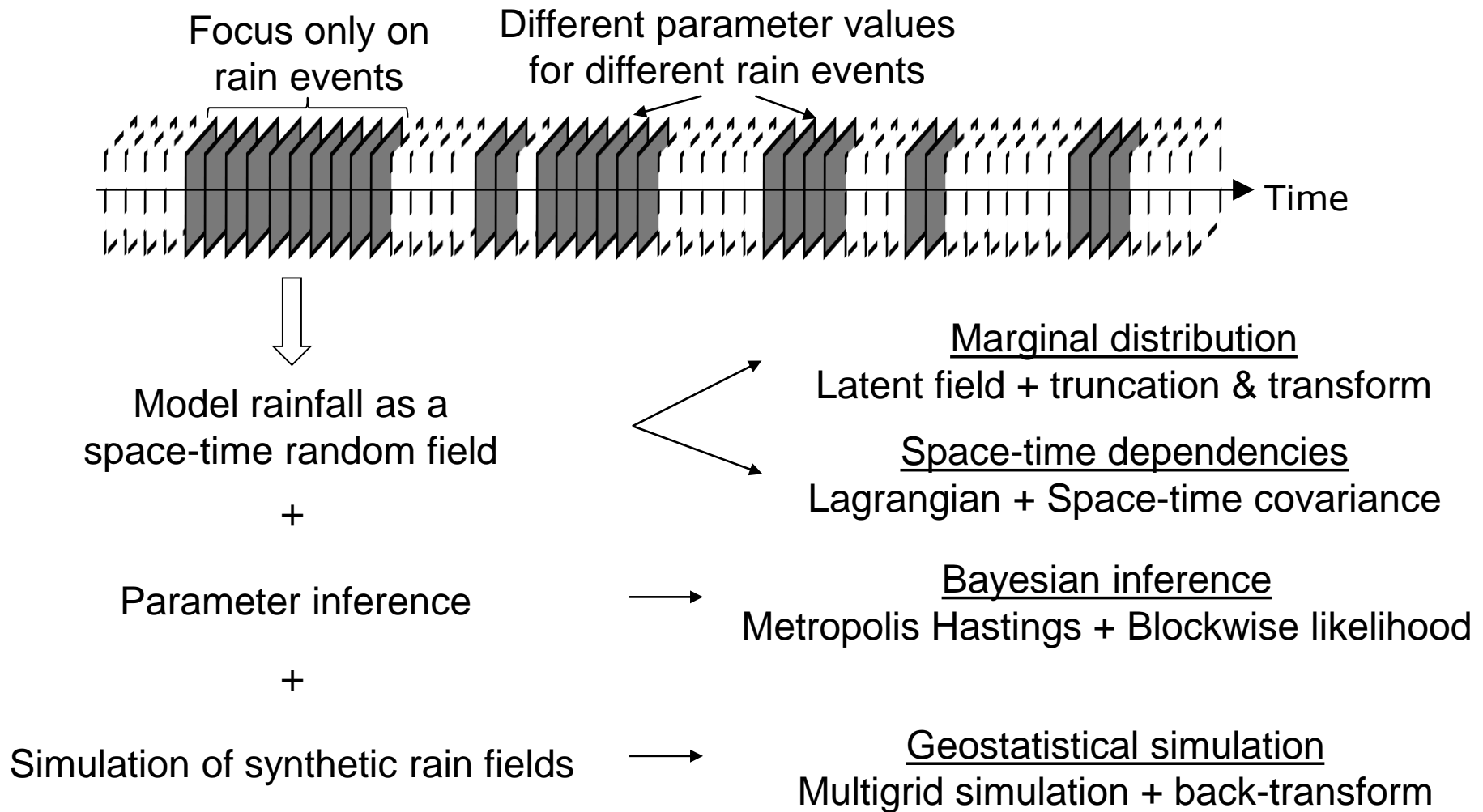
Convective rain

	C1	C2	C3
dt = 30s ds = 50m	MAE = 1.79 Bias = -0.02	MAE = 2.26 Bias = 0.30	MAE = 2.03 Bias = 0.29
dt = 300s ds = 500m	MAE = 3.83 Bias = 0.16	MAE = 1.99 Bias = 0.20	MAE = 2.28 Bias = 0.35

MAE: Mean Absolute Error

- Almost unbiased prediction.
- Low prediction errors.
- Positive contribution of the complex space-time covariance function.

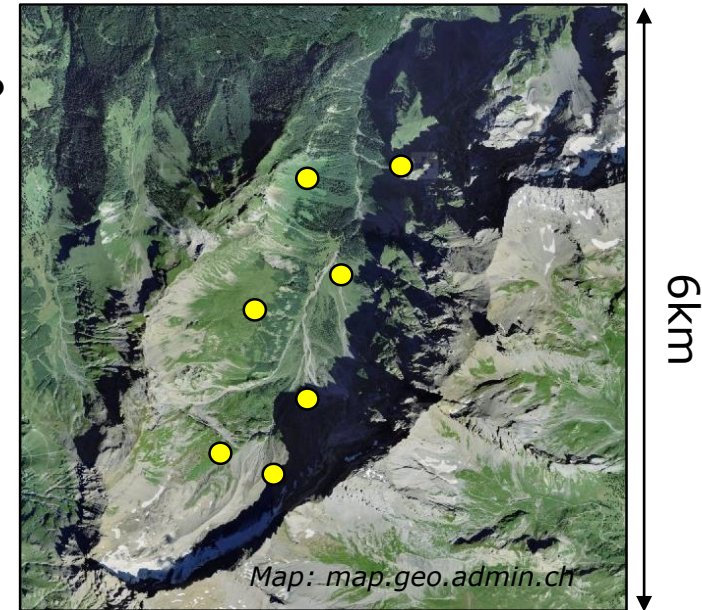
Model summary



Perspectives

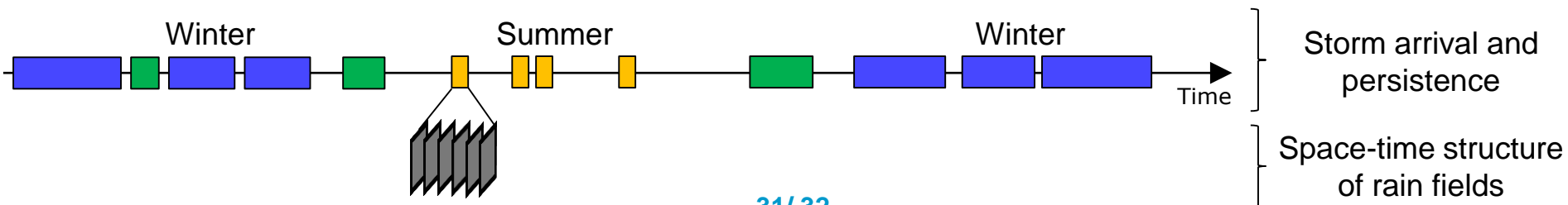
- Apply the proposed framework to a larger network:

- 6km x 6km (Vallon de Nant, Swiss Alps).
- Altitude: 1300 – 2200m => Orographic effects?
- Application to mountain hydrology.



- Model storm arrival and persistence processes:

- Account for seasonality in rainfall structure (rain types).
- Application: stochastic rainfall generation for the whole year.



Conclusion

- Drop counting rain gauges allow to monitor local rain fields:
 - High resolution (0.01mm of rain) => high temporal resolution (30 sec).
 - Low cost, easy to set up => dense networks.
- A local scale stochastic rainfall model has been proposed to handle the features of rainfall arising from HR rain gauge measurements:
 - Rain intermittency and skewed distribution of positive rain rates.
 - Rain advection.
 - Temporal morphing of rain patterns.
- This model can be used to:
 - Investigate space-time dependencies within local rain fields.
 - Estimate HR rain fields over small catchments (Mountain / urban hydrology).

Thank you for your attention

- Drop counting rain gauges allow to monitor local rain fields.
- A Stochastic rainfall model has been proposed to handle these data.
- Applications: - Investigate space-time dependencies within rain fields.
 - Estimate HR rain fields over small catchments.

